

**Scaling and universal power laws  
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**Abstract**

Universal power laws are derived directly from their scaling properties. Though the results are generally valid, we make the discussion herein in connection with the pair distribution of the nearest-neighbouring seismic events in temporal series.

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Recently, a great deal of attention is being given[1]-[6] to the pair distribution  $D(\tau)$  of nearest-neighbouring seisms separated by time  $\tau$ . In general, the probability density of  $N$  serial events denoted by  $i$ , occurring at time  $t_i$ , can be written as

$$\rho(t) = \frac{1}{N} \sum_i \delta(t_i - t) , \quad (1)$$

and, similarly, the pair distribution of nearest-neighbours separated by time  $\tau$  is given by

$$D(\tau) = \frac{1}{N} \sum_i \delta(t_{i+1} - t_i - \tau) . \quad (2)$$

A seismic event occurring at time  $t_i$  has a magnitude  $M_i$ , and much attention has been given to the effect of the lower-bound cutoff magnitude  $M_c$  upon the form of such distributions. In particular, it was recently claimed[2, 3] that the pair distribution  $D(\tau)$  can be represented as

$$D(\tau) = Rf(R\tau) , \quad (3)$$

where  $R = t_0^{-1}e^{-\beta M_c}$  is the rate of seisms with magnitude greater than  $M_c$ , and  $f$  is a universal function, in the sense that it does not depend on  $M_c$ , at least. The parameters  $t_0^{-1}$  (which may be viewed as the seismicity rate) and  $\beta$  in the cutoff parameter  $R$  given above are well documented from various statistical analysis of earthquakes,[7]-[9] their particular values being, however, immaterial for the present discussion.

It is easy to see that the pair distribution given by (3) satisfies the scaling relationship

$$\widetilde{D}(\tau) = e^{-\beta M'_c} f(e^{-\beta M'_c} \tau) = pD(p\tau) , \quad (4)$$

where  $\widetilde{D}(\tau)$  is the transform of  $D(\tau)$  under the change of the cutoff, as given by parameter  $p = e^{-\beta(M'_c - M_c)}$ . Renormalization group arguments have been put forward recently[1] in order to determine the function  $D(\tau)$ . We show herein that it can be derived directly from the scaling equation (4), whose elementary solutions for its Laplace transform are universal power laws.

Indeed,  $\widetilde{D}(\tau)$  may actually be viewed as a function  $\widetilde{D}(p, \tau)$  of both  $p$  and  $\tau$ , and equation (4) can also be written as

$$p^{-1}\widetilde{D}(p, \tau/p) = D(\tau) . \quad (5)$$

It is more convenient to write the above equation by its Laplace transform as

$$\widetilde{D}(p, ps) = D(s) . \quad (6)$$

According to equation (6) the derivative of  $\widetilde{D}$  with respect to  $p$  must vanish. Making use of the new variable  $u = ps$  we obtain

$$p\partial\widetilde{D}(p, u)/\partial p + u\partial\widetilde{D}(p, u)/\partial u = 0 . \quad (7)$$

By separation of variables we get the elementary solution

$$\widetilde{D}(p, u) = p^\beta u^{-\beta} , \quad (8)$$

or, by (6),

$$D(s) = s^{-\beta} , \quad (9)$$

where  $\beta$  is a real exponent. This is a universal power law. It corresponds to a power-law function  $D(\tau) \sim 1/\tau^{1-\beta}$  for  $\beta > 0$ .

In general, the solution of equation (7) is a superposition of elementary solutions of type  $s^{-\beta}$ . Such a particular superposition for integer  $\beta < 0$  (which is related to the hypergeometric function) may give  $D(s) = (1 + s)^{-\alpha}$ , which corresponds to  $D(\tau) \sim (1/\tau^{1-\alpha})e^{-\tau}$ . The normalizing constant for this probability distribution is  $1/\Gamma(\alpha)$ , where  $\Gamma$  is Euler's gamma function. This probability distribution corresponds to  $D(\tau) = RC \cdot [1/(R\tau)^{1-\alpha}]e^{-R\tau/B}$  discussed in Refs.1-3, where  $B = 1.58$ ,  $C = 1/2$  and  $\alpha = 0.66$ , such that  $CB^\alpha\Gamma(\alpha) = 1$ . The exponential form for large values of  $\tau$  is in fact expected from the uncorrelated seismicity,[10] while the power law for small values of  $\tau$  indicates a clustering process.[11]

Apart from the scaling law (3),[4] another main issue in this connection[1]-[4] resides in the particular values assumed by the exponent  $\alpha$  and constant  $B$  in the universal function  $f(\tau) = (1/\tau^{1-\alpha})e^{-\tau/B}$ . It was shown[12] that seismic energy  $E$  can be accumulated in a seismic focus over time  $t$ , such that  $t \sim E^r$ , where  $r$  is a parameter related to the geometry of the focal zone and the focal mechanism. For a uniform mechanism of accumulating energy in a localized, point-like seismic focus the parameter  $r$  acquires the value  $r = 1/3$ . This parameter contributes to the exponent  $\beta = br$  in the magnitude distribution of the earthquakes, for instance in the exceedence rate  $N_{ex}/T = t_0^{-1}e^{-\beta M}$  of earthquakes with magnitude greater than  $M$  (recurrence law), where  $t_0^{-1}$  is the seismicity rate over period  $T$  and  $b = 3.5$  is the Gutenberg-Richter coefficient in the energy-magnitude relationship. For  $M = M_c$  this exceedence rate is the cutoff parameter  $R$ . If energy  $E$  is released in time  $\tau$ , then its rate is given by  $E \sim 1/\tau$  for short times  $\tau$ , according to Omori's law,[13]-[15] so that the above relationship implies  $t \sim E^r \sim 1/\tau^r$ . On the other hand, it is natural to assume that the pair distribution is proportional to time  $t$ , so we are led to conjecture the form

$$D(\tau) \sim \frac{1}{\tau^r} \quad (10)$$

for the pair distribution for small values of  $\tau$ , which implies the exponent parameter  $\alpha = 1 - r$  in the universal function  $f(\tau)$ . For  $r = 1/3$  ( $\beta = 1.17$ ) we get  $\alpha = 2/3$  in agreement with the value indicated in Refs. 1-3.

The exceedence rate given above is well documented for a variety of regions, time intervals and magnitudes. For instance, a worldwide analysis for  $5.8 < M < 7.3$  indicates  $r = 0.39$  ( $\beta = 1.38$ ;

and  $-\ln t_0 = 12.65$  for  $t_0$  measured in years).[16] Data for Southern California[5] seem to indicate  $r = 0.66$  ( $\beta = 2.3$ ; and  $-\ln t_0 = 17.25$  for  $t_0$  in years). Similarly, a recent analysis for Vrancea earthquakes[17] indicates  $r = 0.54$  ( $\beta = 1.89$ ; and  $-\ln t_0 = 9.68$  for  $t_0$  in years). This rather limited variability in the exponents of the pair distribution for small values of  $\tau$  seems to be supported by data.[2]-[4] In particular, the analysis of the pair distribution for 1999 earthquakes with magnitude  $M > 3$  recorded in Vrancea between 1974 and 2004 indicates an exponent  $\alpha = 0.75$ , corresponding to  $r = 0.25$ , in fair agreement with the exponent  $r = 0.33$ [1]-[3]. The exponent  $r$  in (10) exhibits a tendency toward higher values in the limit  $R\tau \rightarrow 0$ , where Omori's law with exponent unity is more effective. For instance, Vrancea data indicate  $r = 0.57$  in fair agreement with the value of this exponent obtained from fitting the recurrence law ( $r = 0.54$ ).[17]

The origin of the parameter  $B$  in the universal function  $f(\tau)$  seems to reside in the fact that for large values of  $\tau$  (where this parameter is effective), *i.e.* for small values of  $t$  (and  $E$ ), the actual relationship between  $t$  and  $E$  is  $1 + t/t_0 = (1 + E/E_0)^r$ , which gives a recurrence law  $N_{ex}/T = t_0^{-1}(1 + e^{bM})^{-r}$  (where  $E_0$  is the threshold energy in the Gutenberg-Richter relationship).[12] This modified recurrence law implies an apparent decrease in the cutoff parameter  $R$  from  $R = t_0^{-1}e^{-\beta M_c}$  to  $R = t_0^{-1}(1 + e^{bM})^{-r}$  (where  $\beta = br$ ), which amounts, for  $M = 0$ , for instance, to a correction factor  $\sim 1.26$ . It is this parameter  $B$  which may account for such correction factors. For instance, the data analysis for Vrancea indicates  $B = 1.17$  ( $C = 0.71$ ), which indeed compares well with  $2^r = 1.19$  corresponding to  $r = 0.25$ . In general, it is worth noting that exponent  $1 - \alpha = r$  in the class of functions  $D(\tau) = RC \cdot [1/(R\tau)^{1-\alpha}]e^{-R\tau/B}$  is a fitting parameter.

In conclusion, the scaling equation (3) indicates the general form of the universal function of the pair distribution for the nearest-neighbouring earthquakes, which consists of a certain superposition of elementary power laws, while the parameters in this universal function, namely the exponent parameter  $\alpha$  and parameter  $B$ , are related to particular seismicities, exhibiting a relatively limited variability, such that it may be said that the pair distribution is rather represented by a quasi-universal function.

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