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The best prediction is about the present

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The magnitude M of some (correlated) foreshocks, occurring at time t before the occurrence time t_{ms} of a mainshock with magnitude M_0 , is

$$M = \frac{1}{b} \ln \frac{t_{ms} - t}{\tau_0} , \quad (1)$$

where $b = 3.45$ is the Hanks-Kanamori constant, the small threshold time τ_0 is given by

$$\tau_0 = r t_0 e^{-b(1-r)M_0} , \quad (2)$$

and r and t_0 are the parameters of the background seismicity; for Vrancea $r = 0.65$ and $-\ln t_0 = 11.32$ years; the time t varies from some value smaller than t_{ms} ($M < M_0$) up to $t_{ms} - \tau_0$. By fitting equation (1) to a set of correlated foreshocks, we get the fitting parameters t_{ms} and τ_0 , such that we are able to predict the occurrence time and the magnitude of the mainshock.

The function $M(t)$ given by equation (1) has an abrupt variation for t near t_{ms} . The errors are

$$\delta M = \frac{\delta t_{ms}}{b(t_{ms} - t)} + (1 - r)\delta M_0 . \quad (3)$$

These errors are very large for t close to t_{ms} . In order to minimize the errors, and to get a "more exact" parameter t_{ms} (small δt_{ms}), we need a good fit, *i.e.* many data near t_{ms} . These would be microtremors (*i.e.*, earthquakes with small magnitudes), as Omori's law shows indeed for accompanying seismic events. When these microtremors occur, we are already at the moment t_{ms} , which we want to predict (the so-called nucleation phase of the mainshock); while performing the fit, we are already experiencing an ongoing mainshock. Our prediction is already fulfilled: it is precisely that "present" moment of time. The best prediction is precisely the realization of the event. The true science is living it.